



EPIC Near-UV Aerosol Product: Status and Validation



Omar Torres

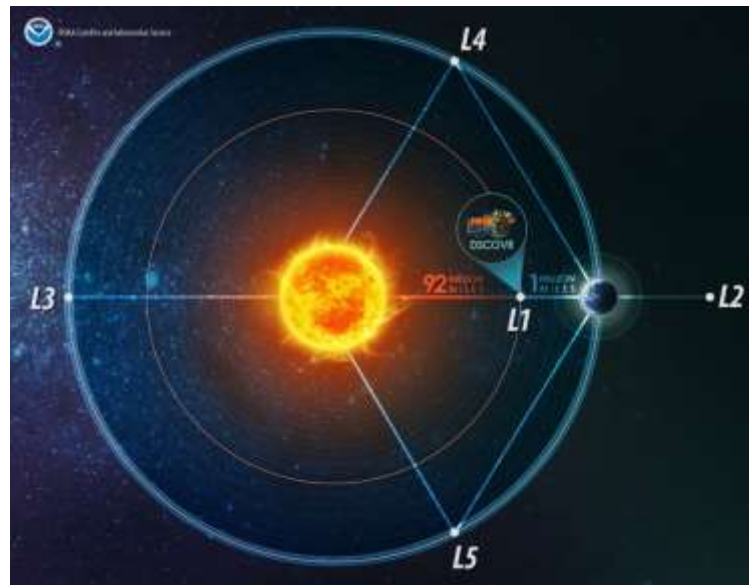
NASA Goddard Space Flight Center

Howard University Beltsville Campus

Fall Series Seminar

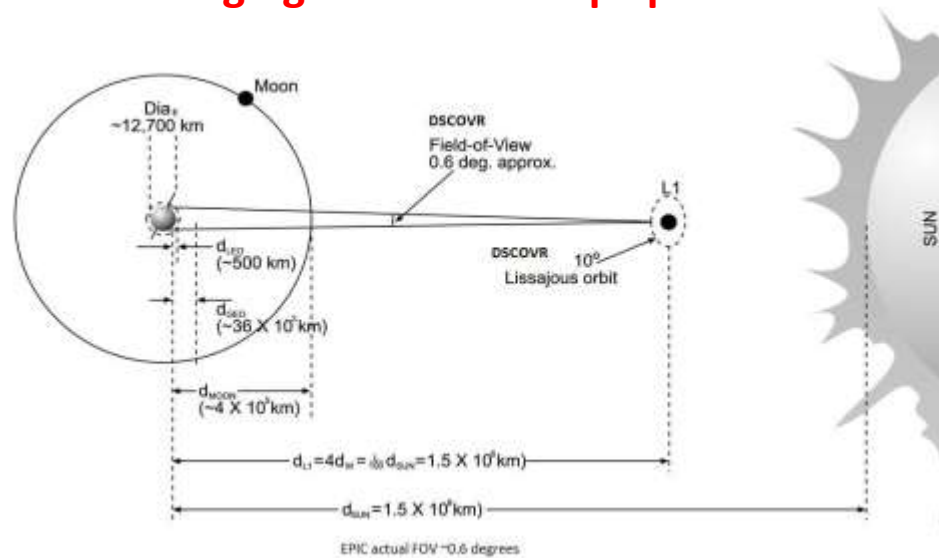
November 22, 2019

DSCOVR location in space



L1 and L2 points are advantageous for Earth remote sensing applications

Earth Panchromatic Imaging Camera – Deep Space Climate Observatory



- EPIC is one of three sensors on the DSCOVR spacecraft.
- Launched on Feb. 11 2015
- Reached L1 point on June 8, 2015
- Observes the Sun-lit face of the Earth from sunrise to sunset every 66 minutes.
- Spatial Resolution ~ 18 km



EPIC Channels and Products

λ (nm)	FWHM (nm)	Nominal Product
<i>317.5 ± 0.1</i>	<i>1 ± 0.2</i>	<i>Ozone</i>
<i>325 ± 0.1</i>	<i>2 ± 0.2</i>	<i>Ozone</i>
<i>340 ± 0.3</i>	<i>3 ± 0.6</i>	<i>Ozone, Aerosols, Clouds</i>
<i>388 ± 0.3</i>	<i>3 ± 0.6</i>	<i>Aerosols, Clouds</i>
<i>443 ± 1</i>	<i>3 ± 0.6</i>	<i>Aerosols</i>
<i>551 ± 1</i>	<i>3 ± 0.6</i>	<i>Aerosols, Vegetation</i>
<i>680 ± 0.2</i>	<i>2 ± 0.4</i>	<i>Aerosol, Vegetation, Clouds, O₂ B-Band Reference</i>
<i>687.75 ± 0.2</i>	<i>0.8 ± 0.2</i>	<i>O₂ B-Band Cloud Height</i>
<i>764.0 ± 0.2</i>	<i>1 ± 0.2</i>	<i>O₂ A-Band Cloud Height, Aerosol Height</i>
<i>779.5 ± 0.3</i>	<i>2 ± 0.4</i>	<i>O₂ A-Band Reference, Vegetation</i>

Moon Transit Movie



EPIC near UV Aerosol Products

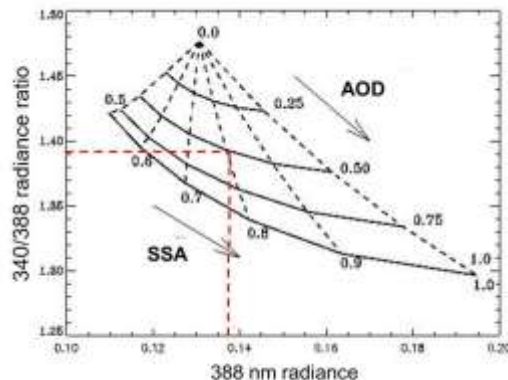
- UV Aerosol Index
- Aerosol Optical Depth (388 nm)
- Single Scattering Albedo (388 nm)

$$UVAI = -100 \left\{ \log \left[\frac{I_{340}^{obs}}{I_{388}^{obs}} \right] - \log \left[\frac{I_{340}^{cal}}{I_{388}^{cal}} \right] \right\}$$

Difference between observed and calculated UV Spectral Contrast

- UVAI is sensitive to AOD, aerosol layer height, Aerosol Absorption Exponent
- UVAI typical values vary from -1 (non-absorbing aerosols) to about 6 (large AOD tropospheric smoke and dust layers)
- Double-digit UVAI values are generally associated with high altitude absorbing aerosol layers (volcanic ash and carbonaceous particles)

Quantitative Products

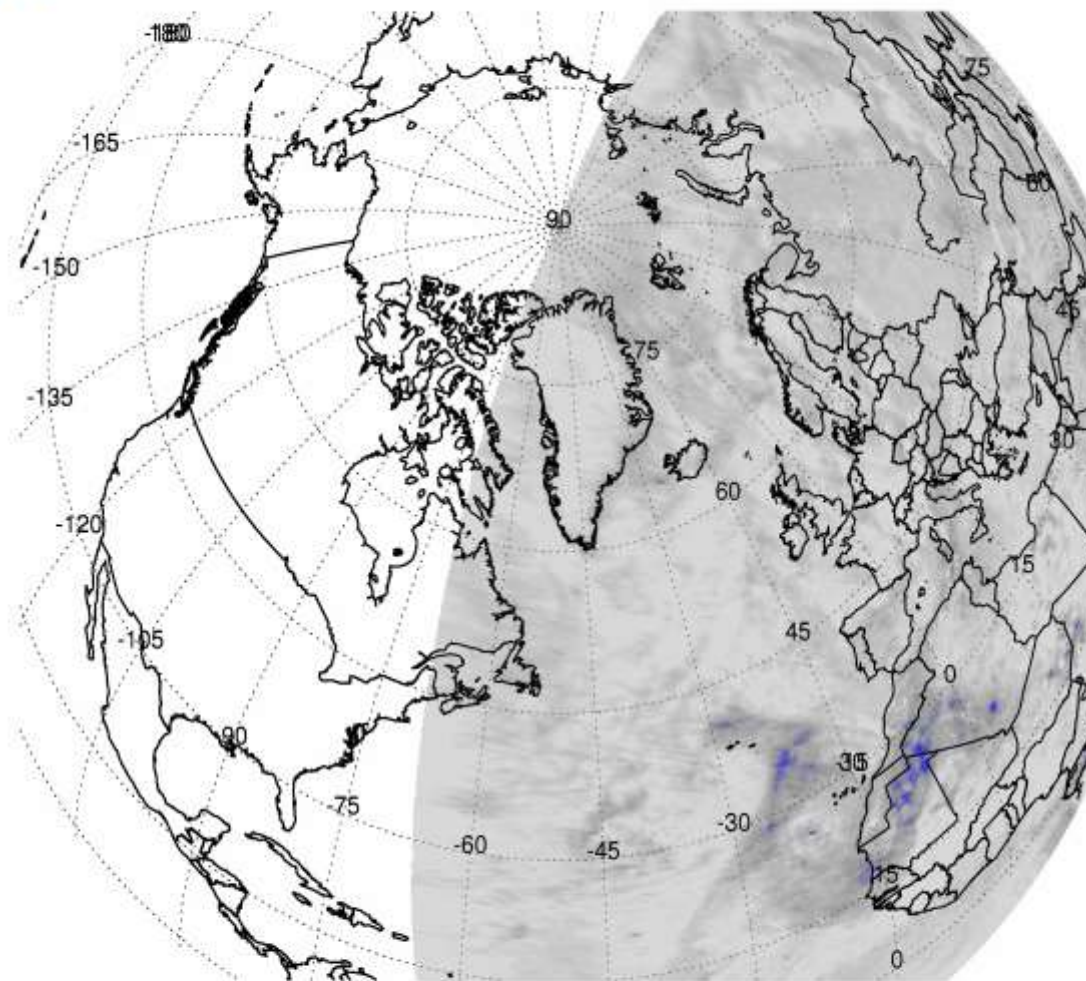


Inversion Scheme:

For a given aerosol type and layer height, satellite measured radiances at 340 and 388 nm are associated with a set of AOD and SSA values.

PyroCb event of August, 2017

EPIC Version 3 L1b data



DSCOVR-EPIC UV Aerosol Index 2017 08 10 10 35



Current Status

- The near-UV aerosol algorithm (EPICAERUV) was applied to the entire record of EPIC Version 2 L1B.
 - The product includes
 - *Aerosol Optical Depth*
 - *Single-scattering Albedo*
 - *Aerosol Absorption Optical Depth*
 - *UV Aerosol Index*
 - *Reflectivity*
 - *Above-cloud aerosol optical depth*
 - *Aerosol-corrected cloud optical depth*
- } **Testing mode**
- Testing of Version 3 L1B on EPICAERUV underway

EPIC near UV Aerosol Products

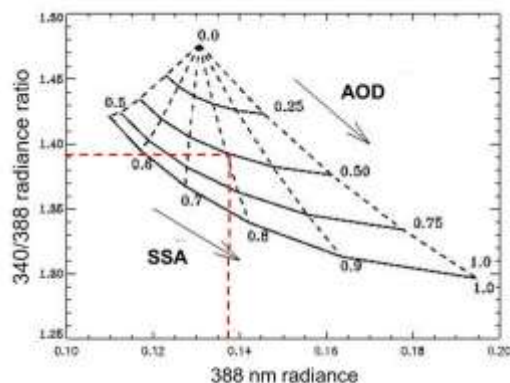
- UV Aerosol Index
- Aerosol Optical Depth (388 nm)
- Single Scattering Albedo (388 nm)

$$UVAI = -100 \left\{ \log \left[\frac{I_{340}^{obs}}{I_{388}^{obs}} \right] - \log \left[\frac{I_{340}^{cal}}{I_{388}^{cal}} \right] \right\}$$

Difference between observed and calculated UV Spectral Contrast

- UVAI is sensitive to AOD, aerosol layer height, Aerosol Absorption Exponent
- UVAI typical values vary from -1 (non-absorbing aerosols) to about 6 (large AOD tropospheric smoke and dust layers)
- Double-digit UVAI values are generally associated with high altitude absorbing aerosol layers (volcanic ash and carbonaceous particles)

Quantitative Products



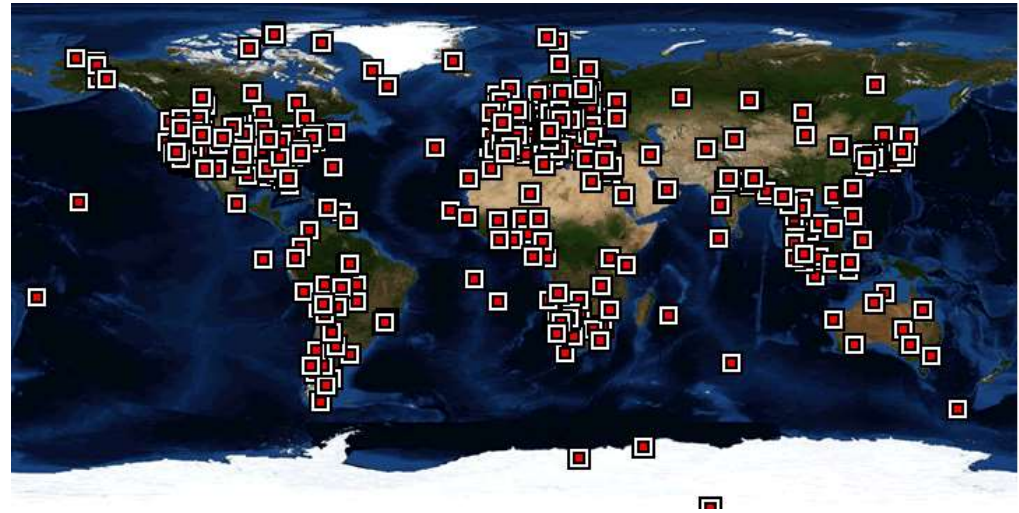
Inversion Scheme:

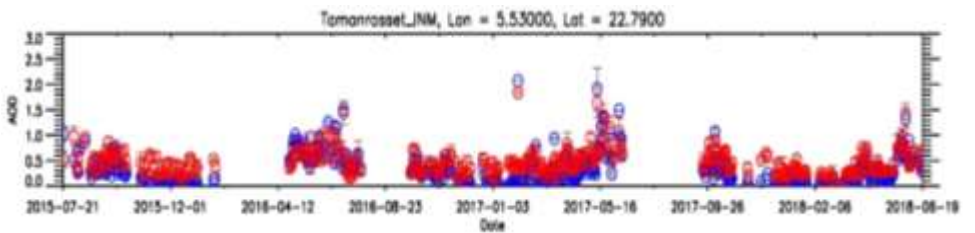
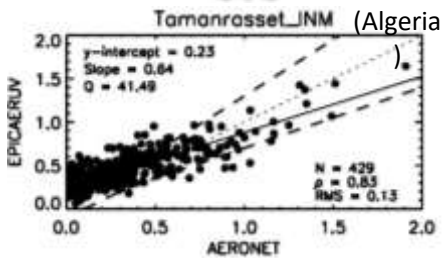
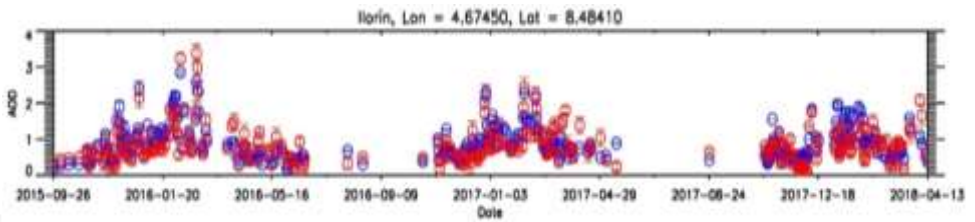
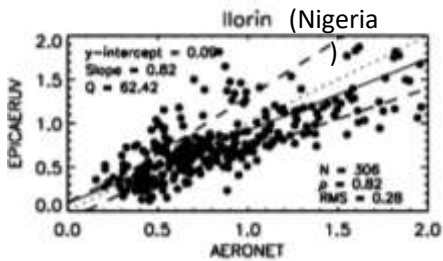
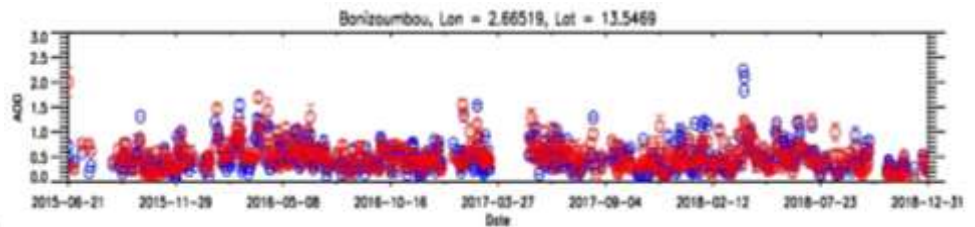
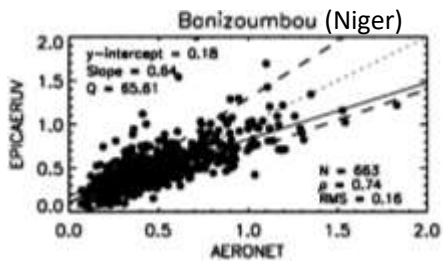
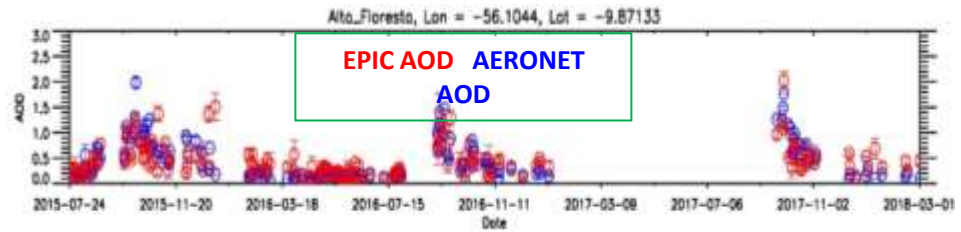
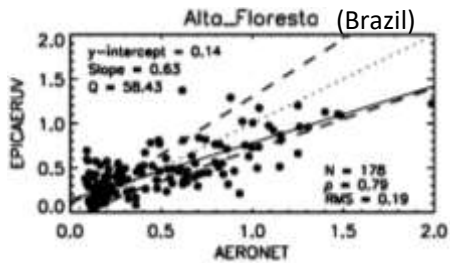
For a given aerosol type and layer height, satellite measured radiances at 340 and 388 nm are associated with a set of AOD and SSA values.

Validation of EPIC Aerosol Product

- Nearly 4-year long EPICAERUV retrievals are validated against ground-based AERONET data set.
- Both AOD and SSA products are evaluated.
- AERONET Version 3 dataset

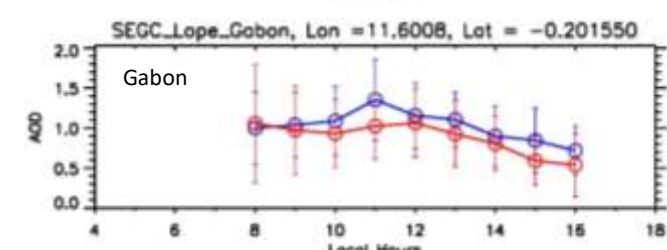
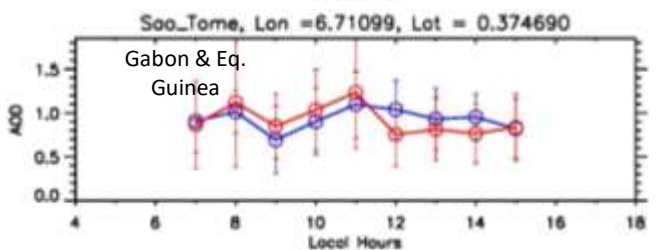
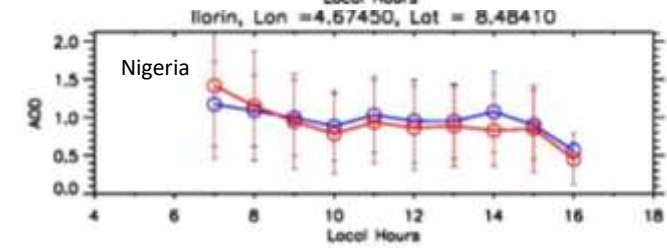
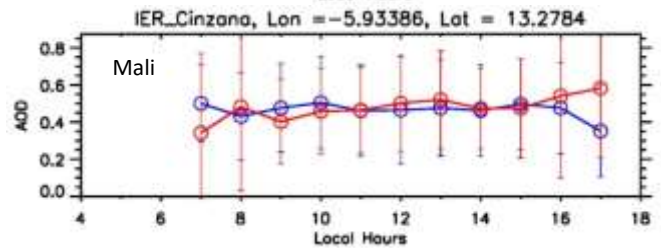
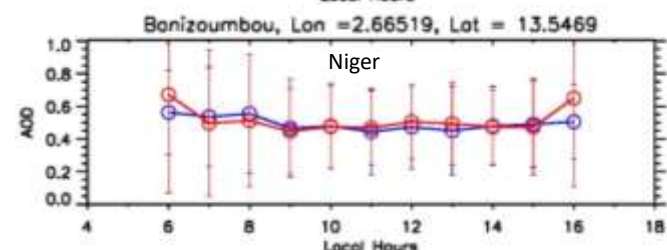
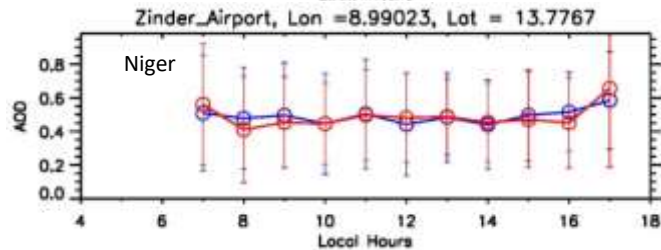
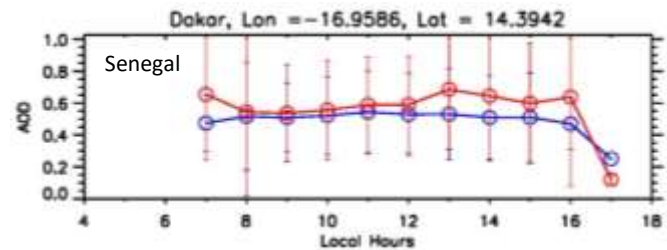
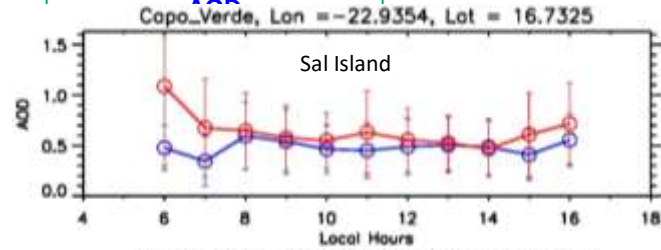
Globally distributed AERONET 2017





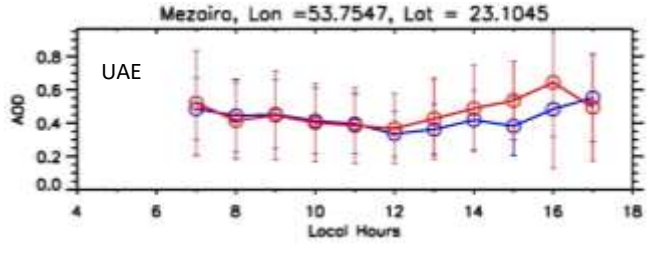
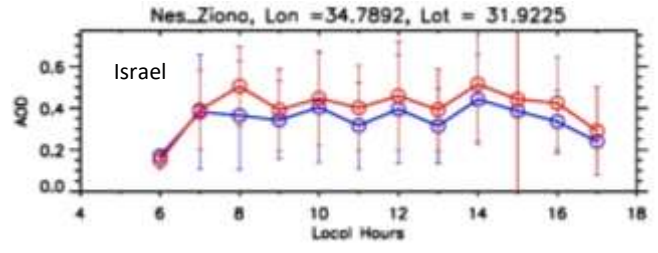
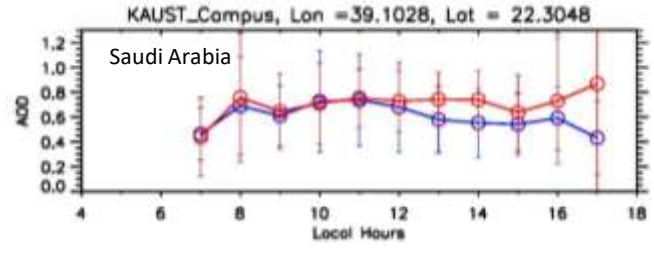
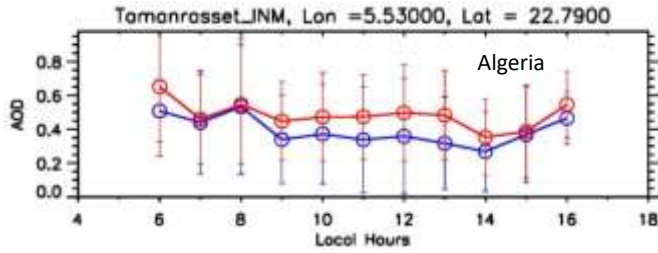
EPIC AOD AERONET

Sahel Region



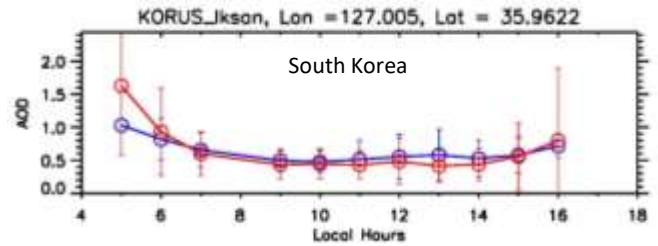
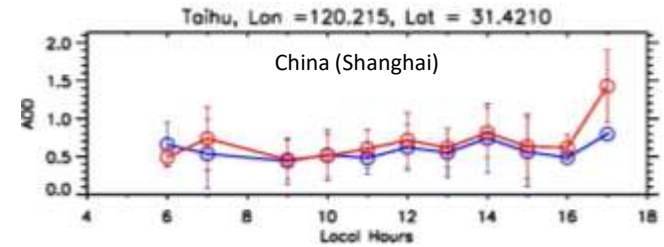
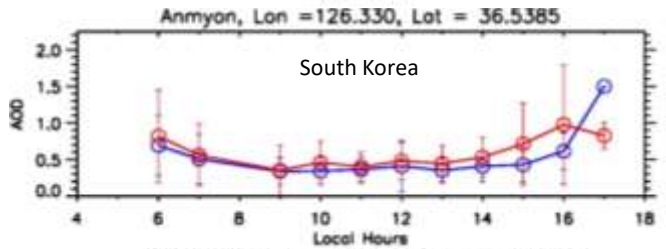
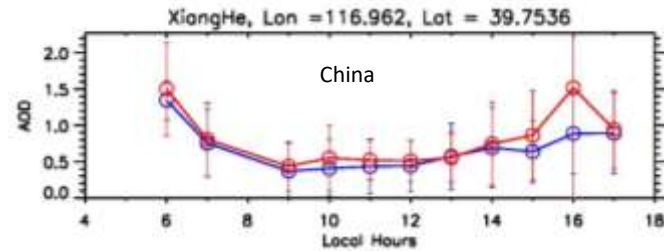
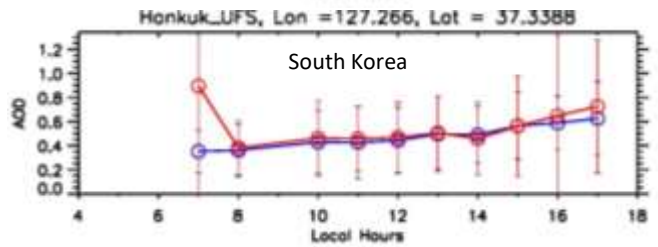
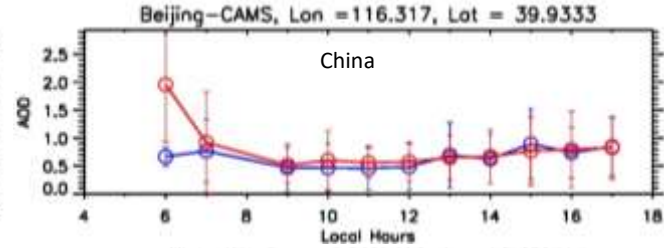
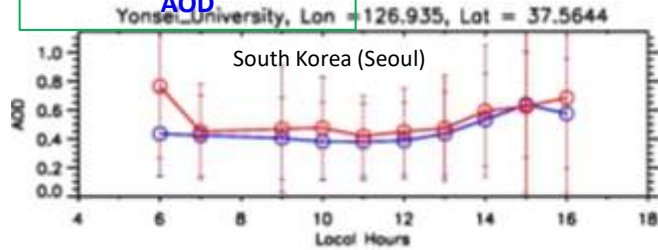
Desert Region

EPIC AOD
AERONET
AOD

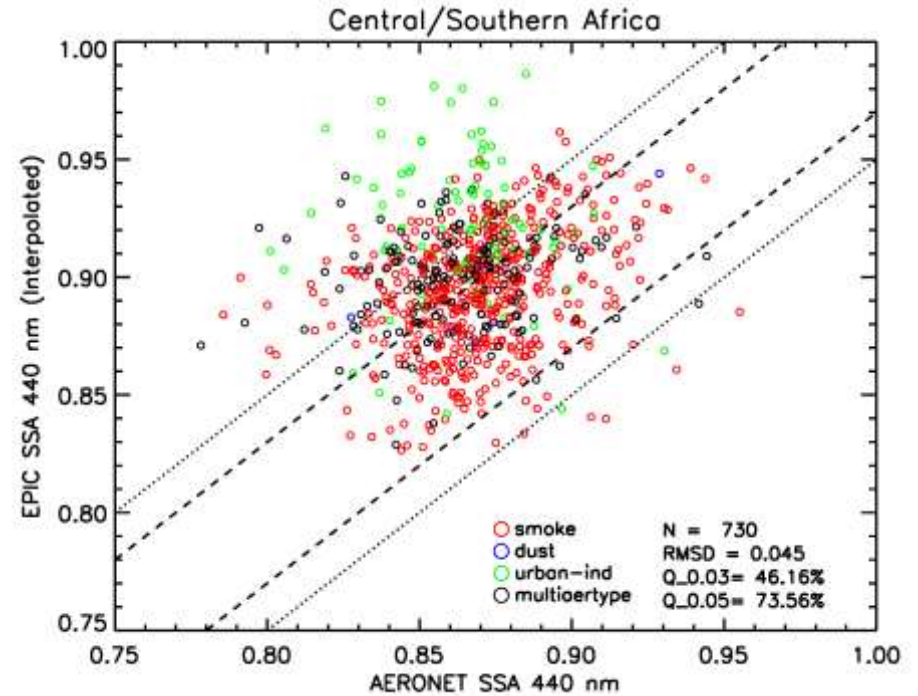
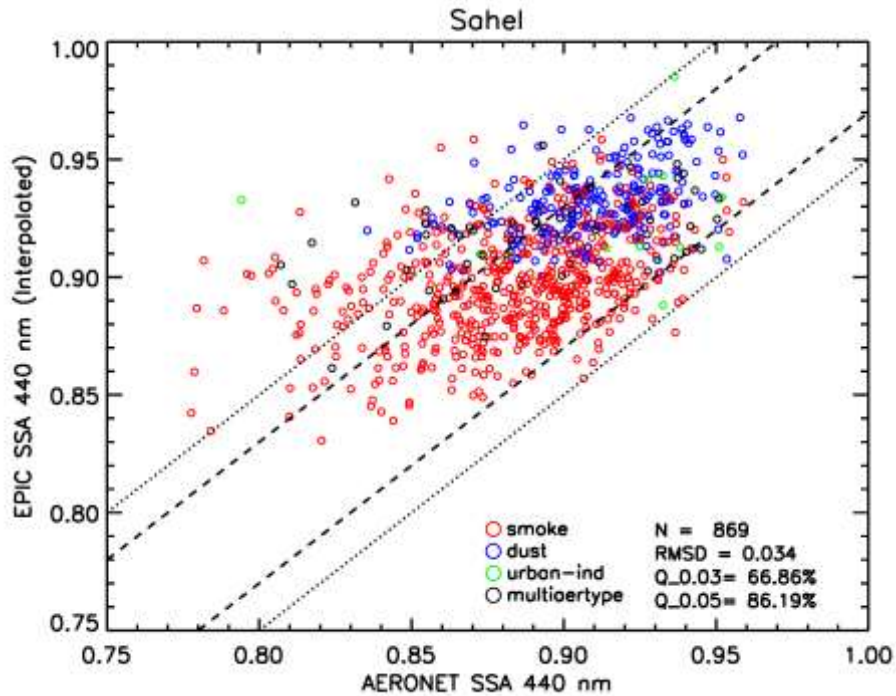


EPIC AOD
AERONET
AOD

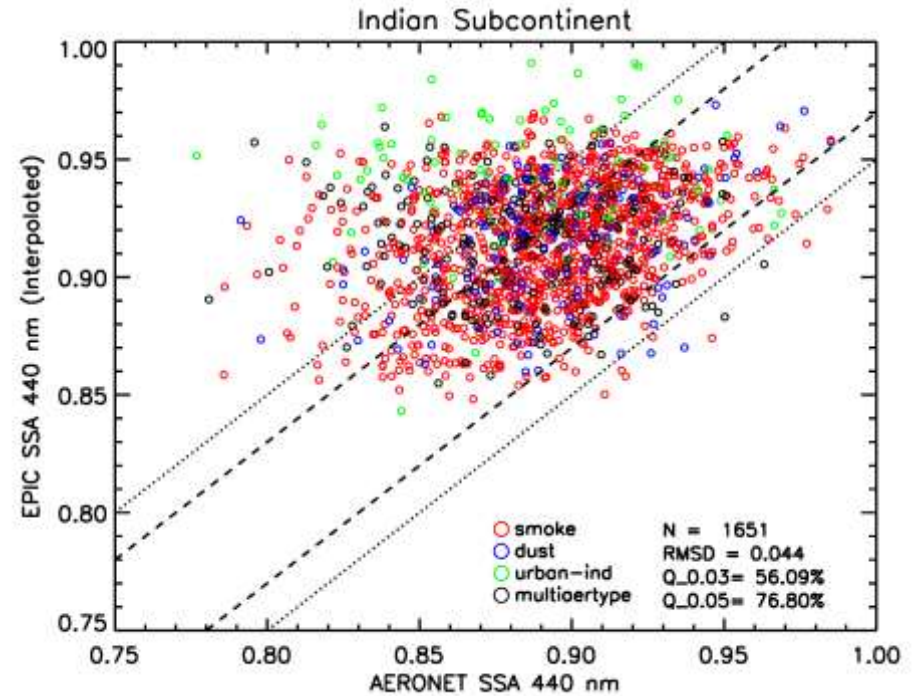
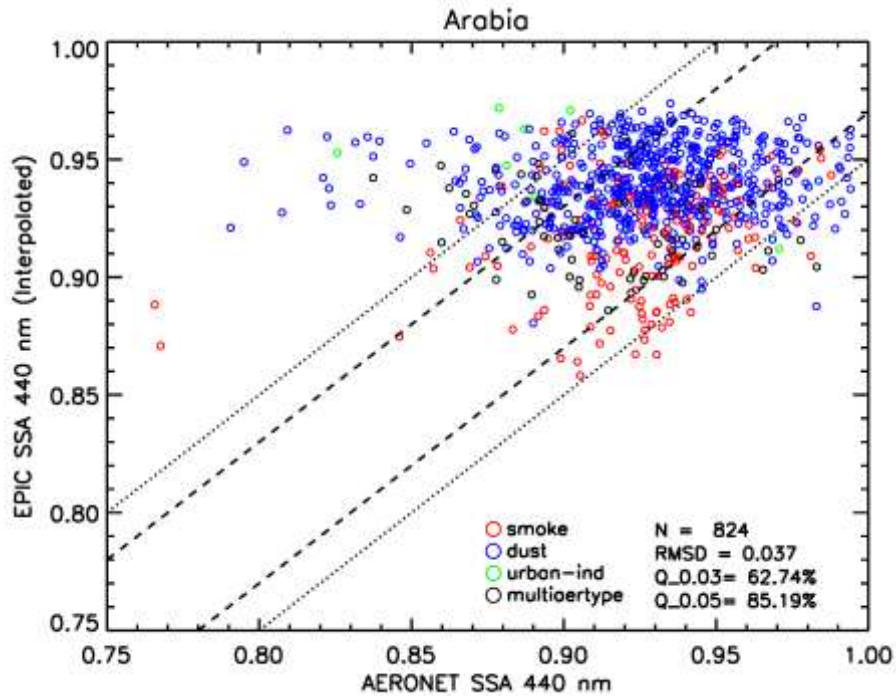
South Korea & China



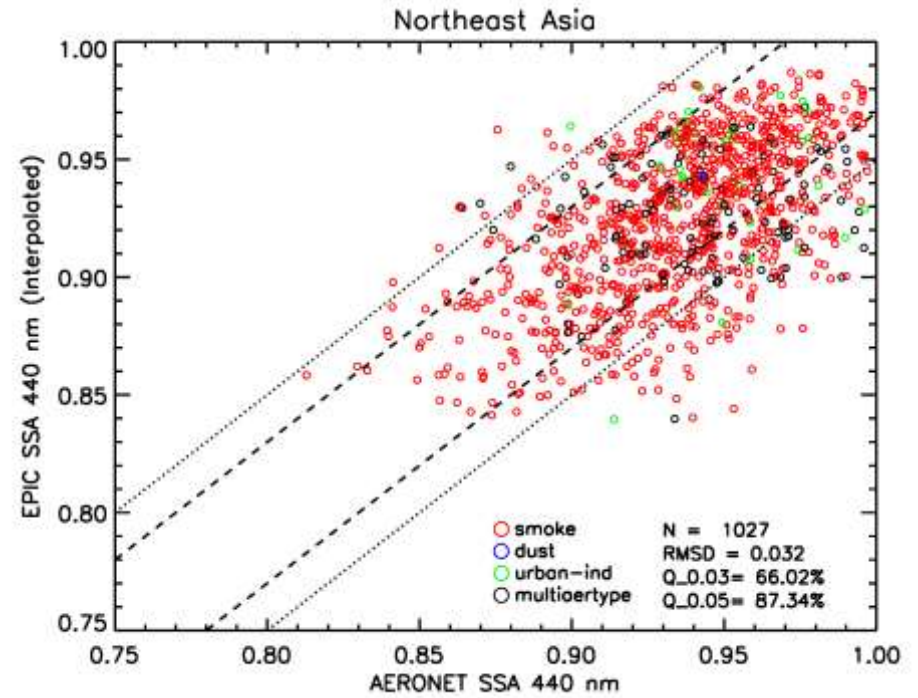
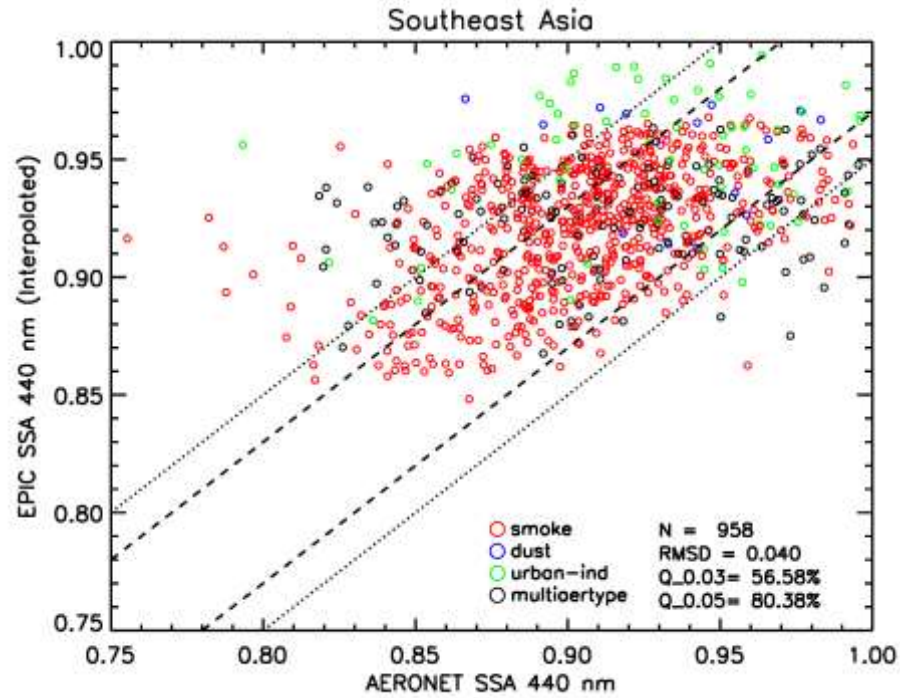
Single Scattering Albedo Evaluation



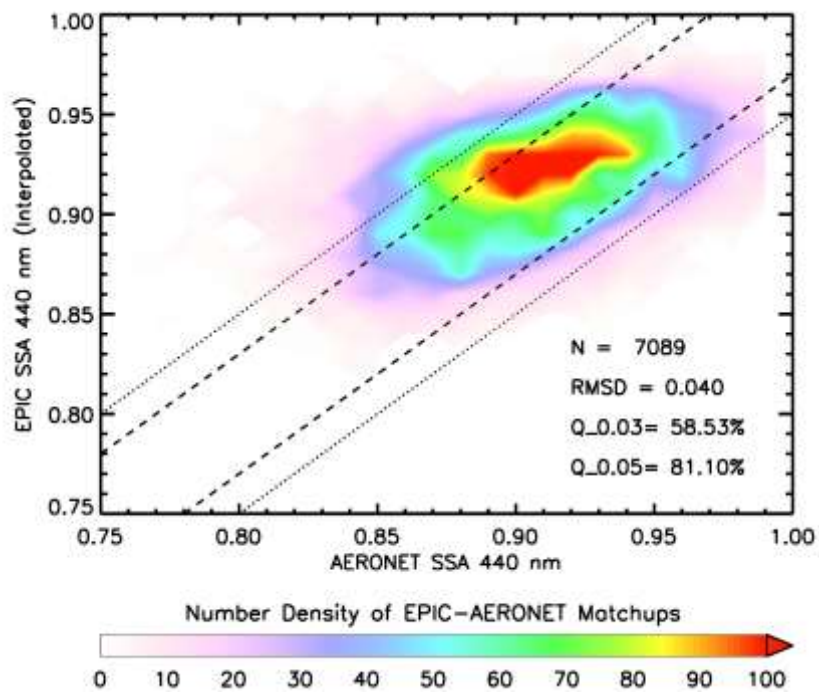
Single Scattering Albedo Evaluation



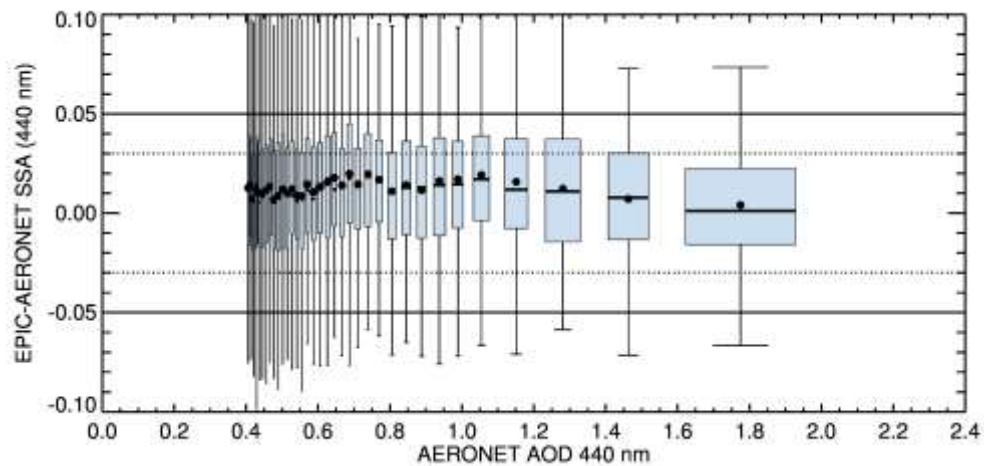
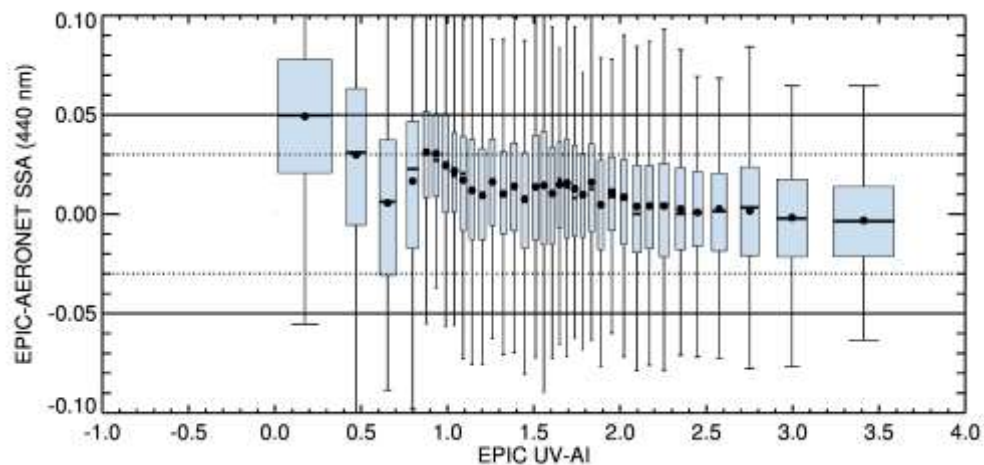
Single Scattering Albedo Evaluation



Summary of Single Scattering Albedo Evaluation



**Nearly 60% and 80% matchups
are within 0.03 and 0.05**



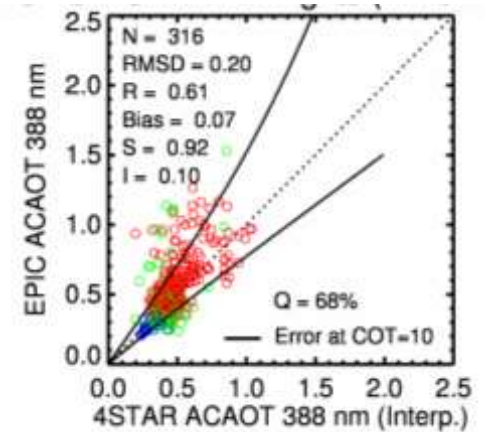
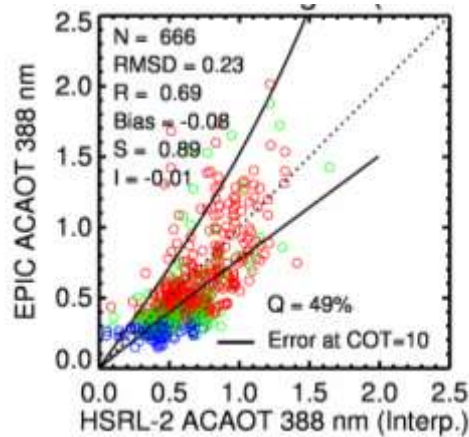
Validation EPIC Above-cloud AOD

Against

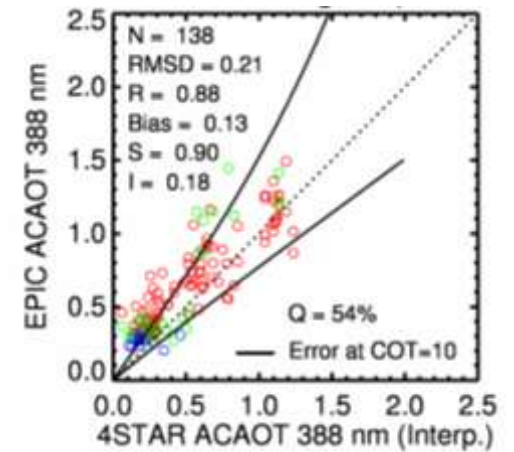
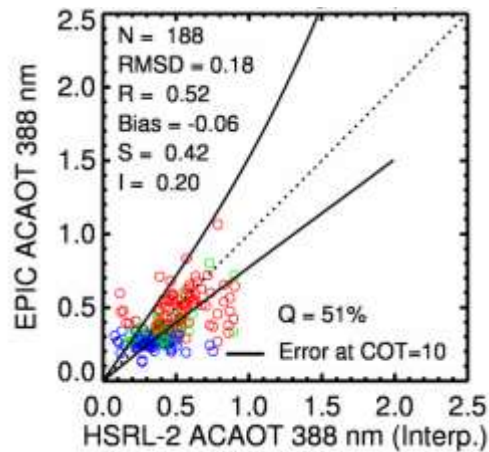
ORACLES airborne
measurements

>50% matchups are within
expected uncertainties

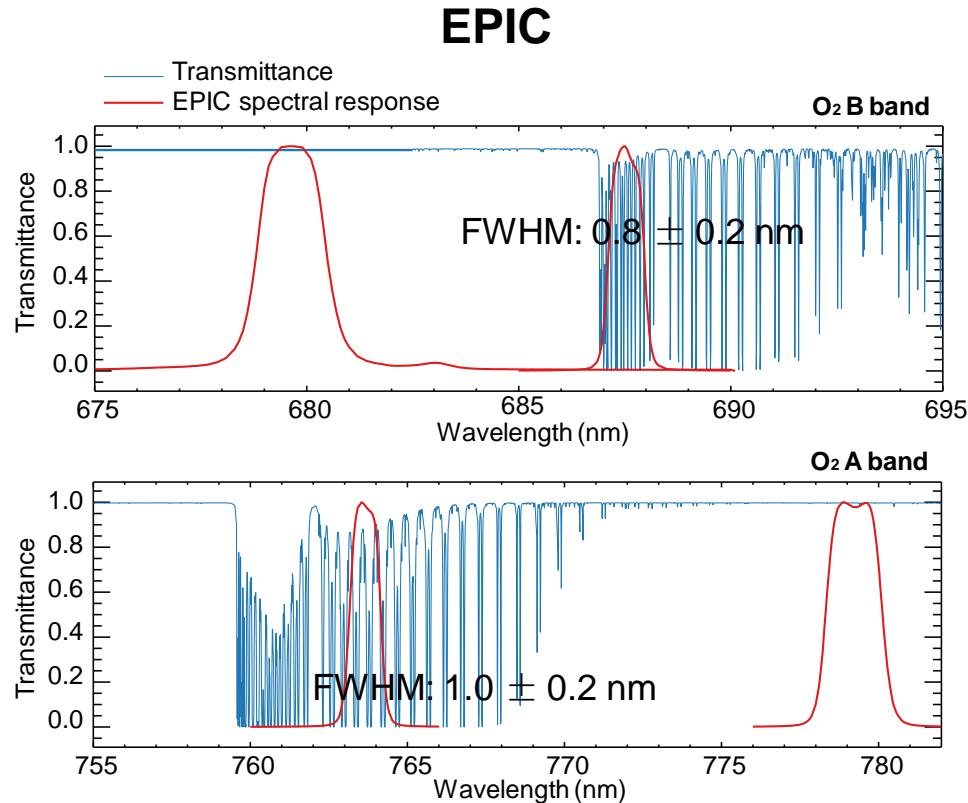
2016



2017

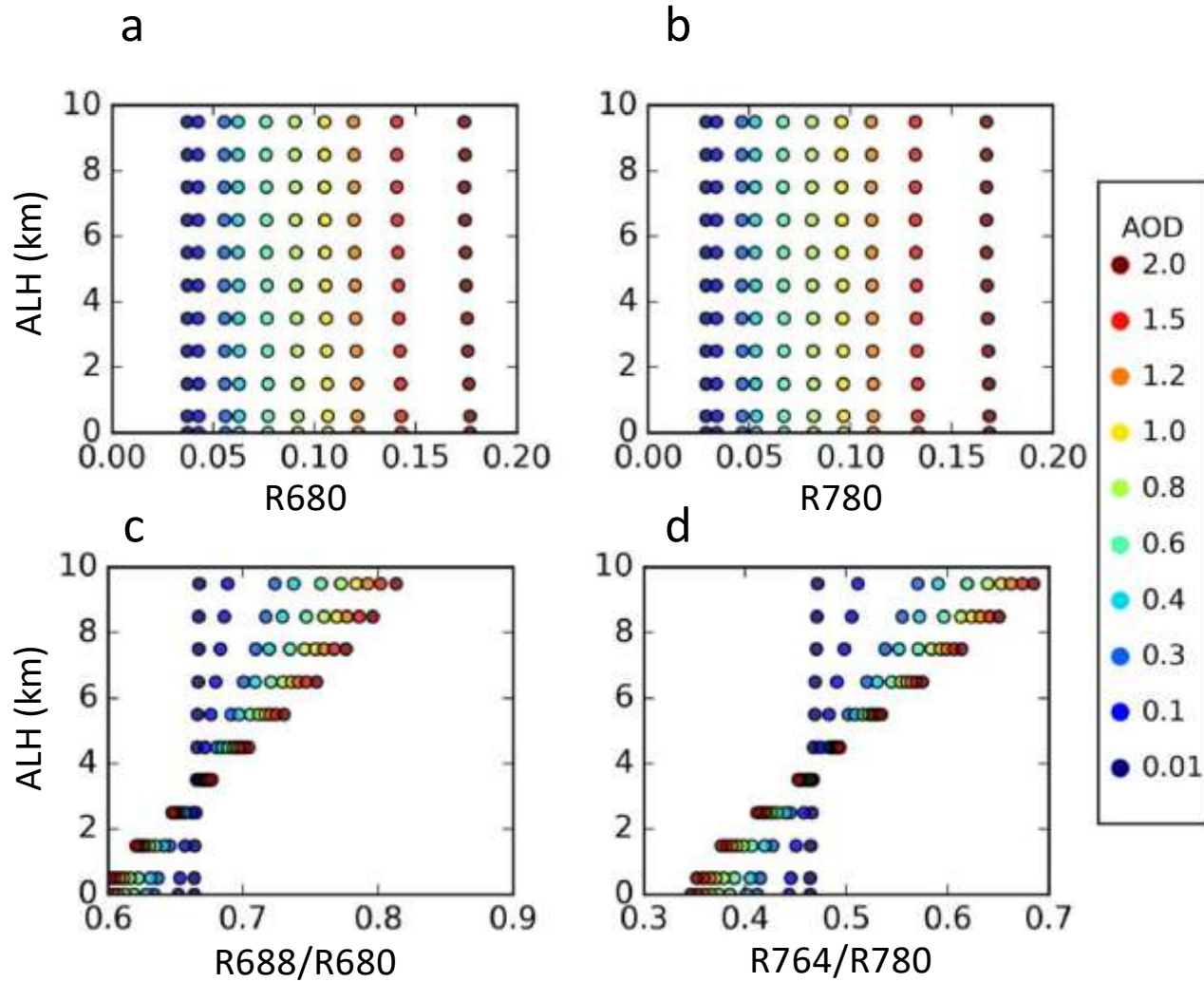


Aerosol Layer Height from O₂ A and B bands of EPIC



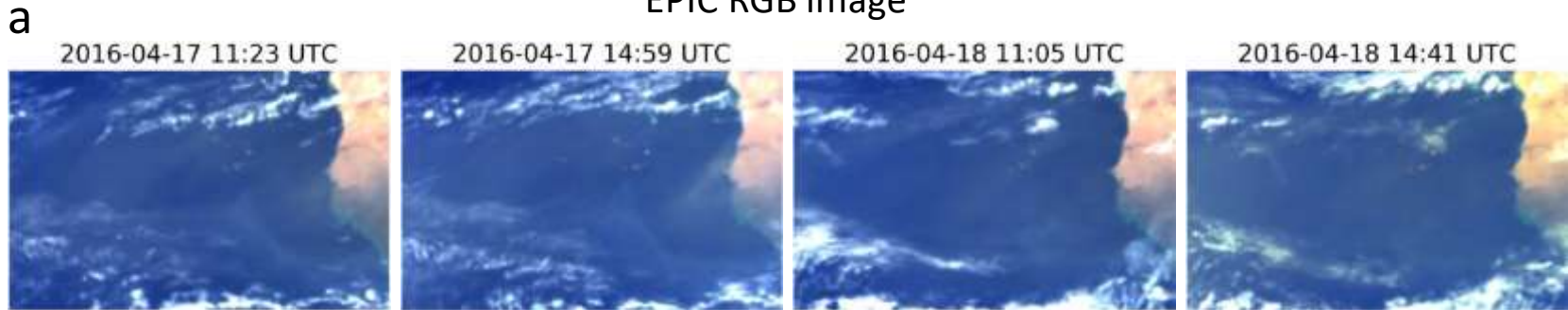
Xu, X., J. Wang, Y. Wang, J. Zeng, O. Torres, Y. Yang, A. Marshak, J. Reid, *and* S. Miller (2017), Passive remote sensing of altitude and optical depth of dust plumes using the oxygen A and B bands: First results from EPIC/DSCOVR at Lagrange-1 point, *Geophys. Res. Lett.*, 44, [doi:10.1002/2017GL073939](https://doi.org/10.1002/2017GL073939).

Simultaneous Retrieval of AOD and ALH

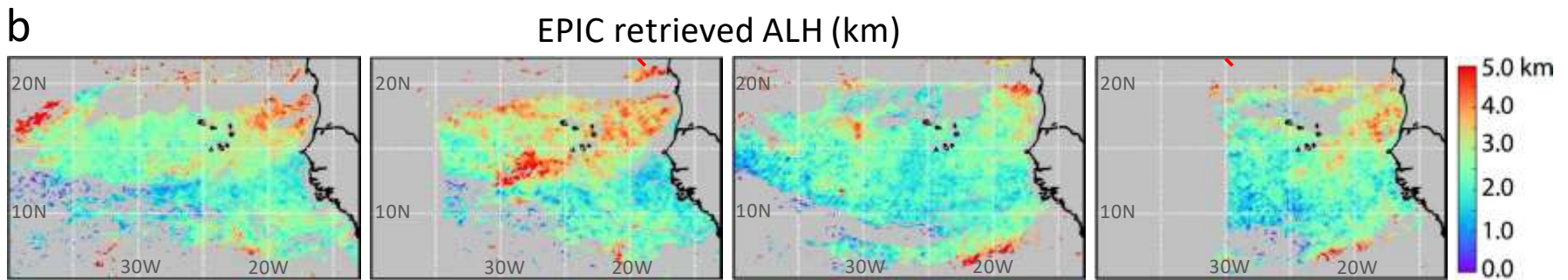


Retrieval of plume height and AOD

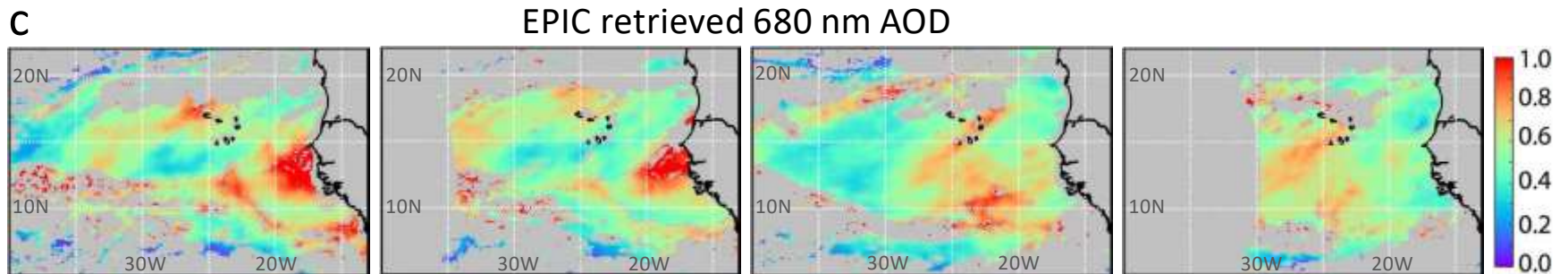
EPIC RGB image



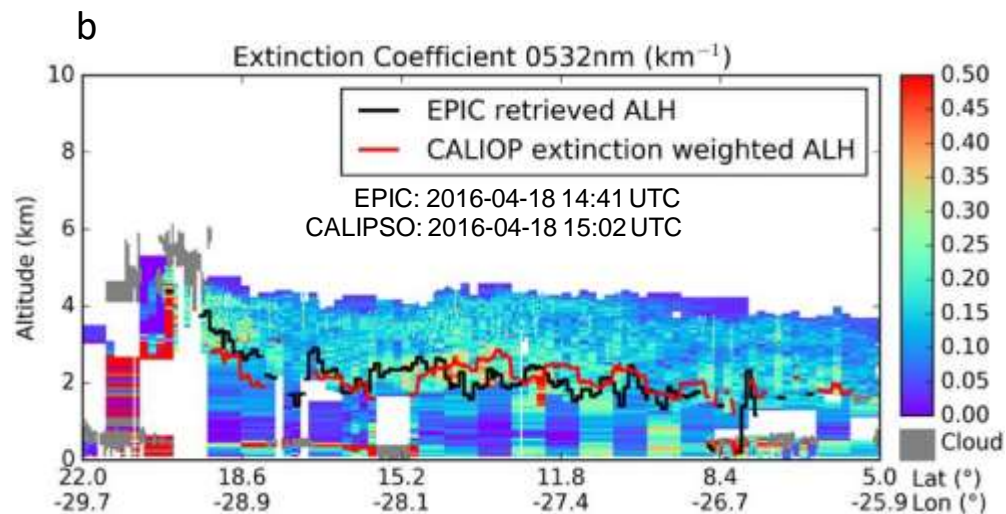
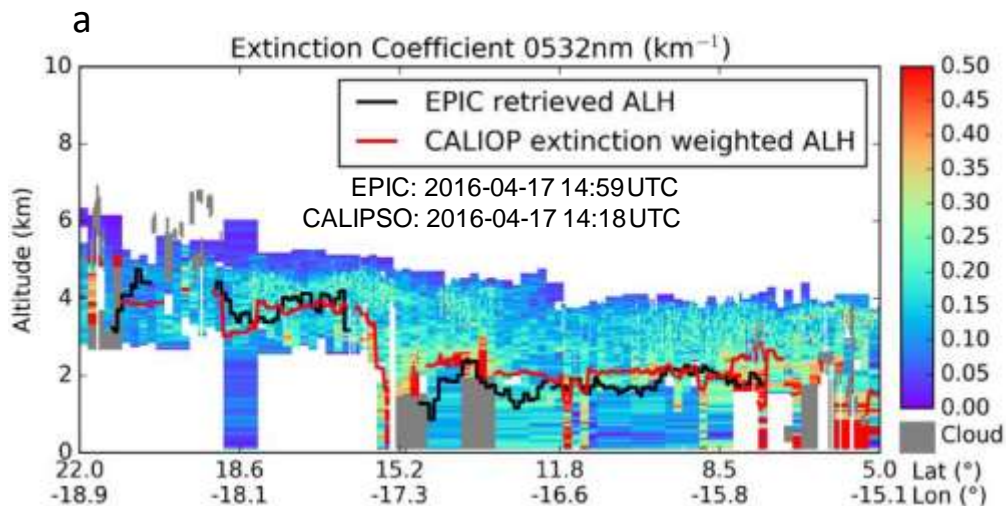
EPIC retrieved ALH (km)



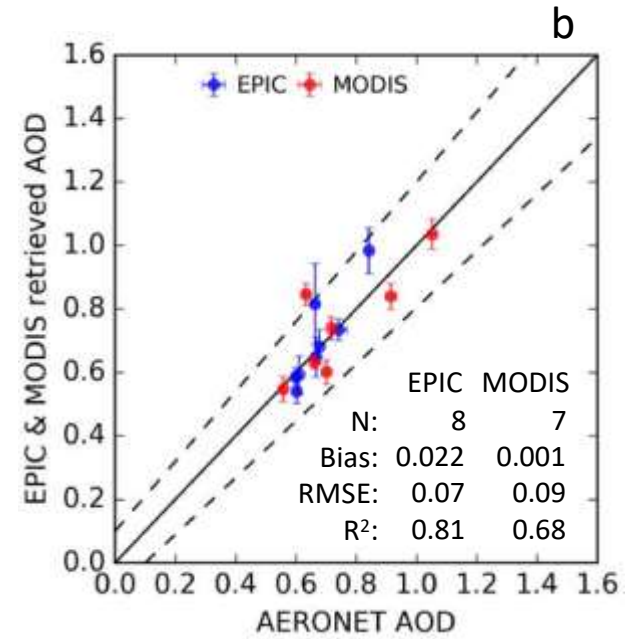
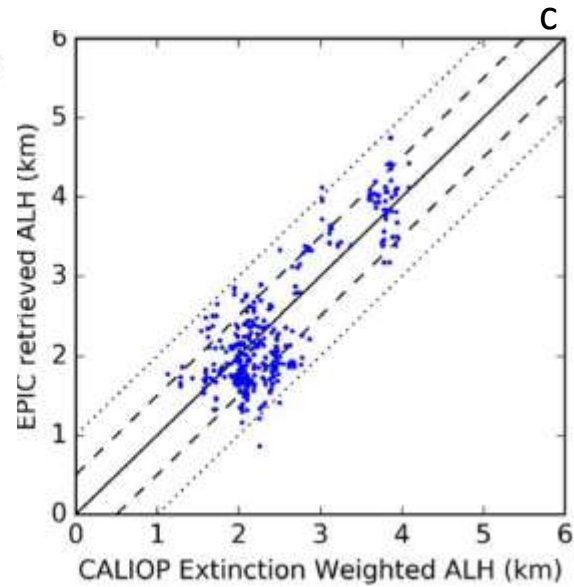
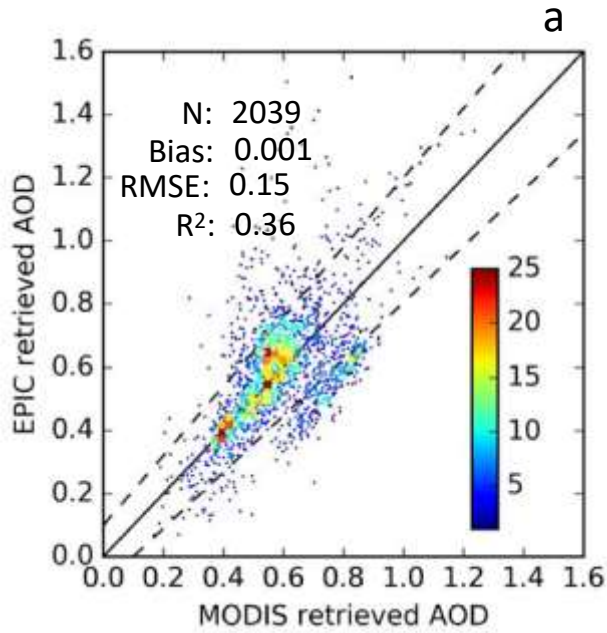
EPIC retrieved 680 nm AOD




Validation with CALIOP and MODIS data



AOD evaluations



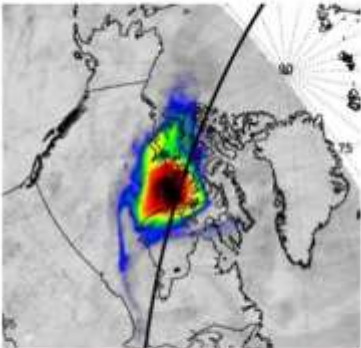
A large, billowing plume of white and grey smoke rises from a forest fire, filling the sky. The smoke is thick and textured, with a dark grey base near the forest and a lighter, almost white top. The forest below is dark green and dense. The sky is a pale, overcast grey.

**Stratospheric Injection of Massive Smoke Plume from Canadian Boreal Fires in 2017:
EPIC verification of self-lifting of carbonaceous aerosol layers**

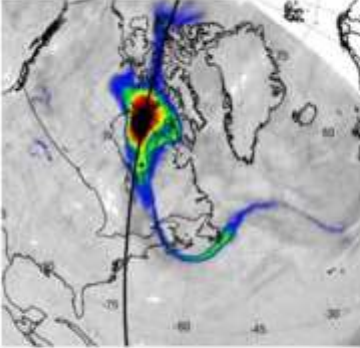
Torres et al., Stratospheric Injection of Massive Smoke Plume from Canadian Boreal Fires in 2017 as seen by DSCOVR-EPIC, CALIOP and OMPS-LP Observations, 2019, JGR (under review)

EPIC and CALIOP view

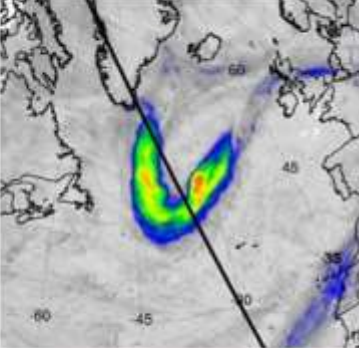
August 15, 2017



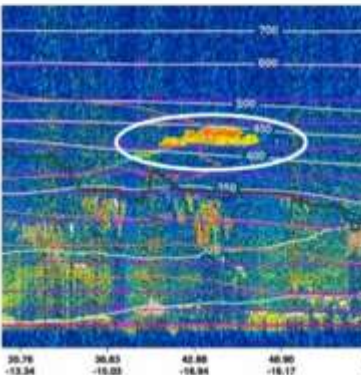
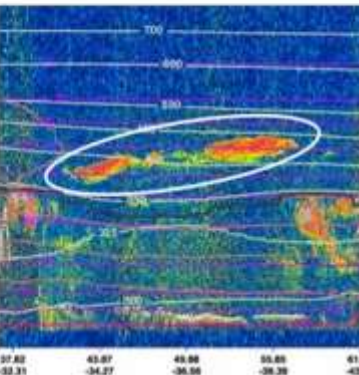
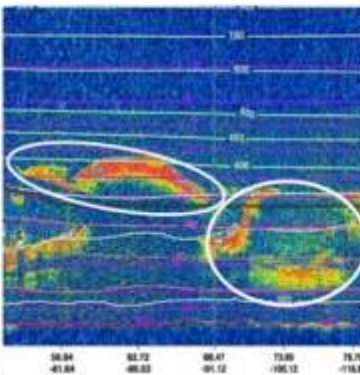
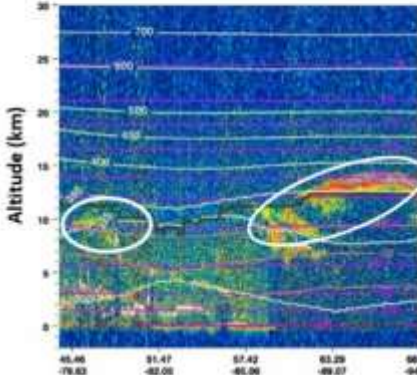
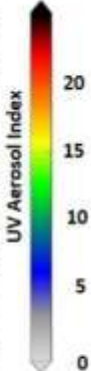
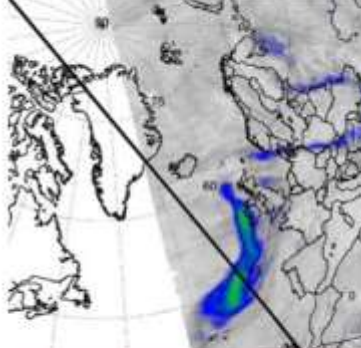
August 17, 2017



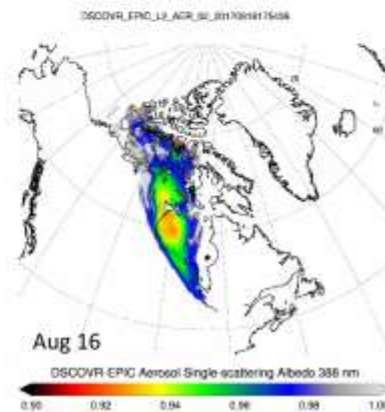
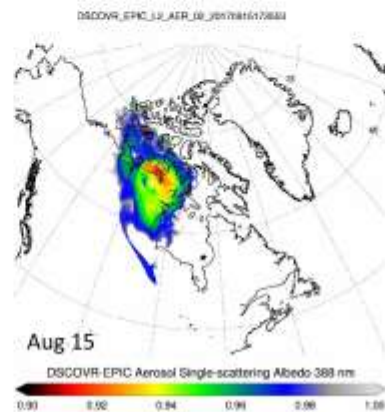
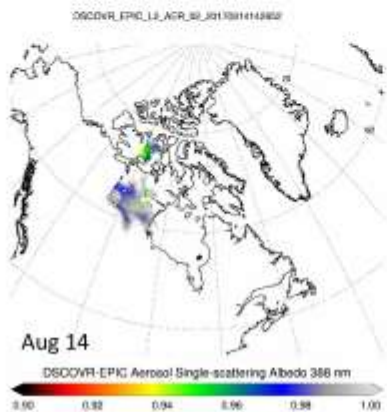
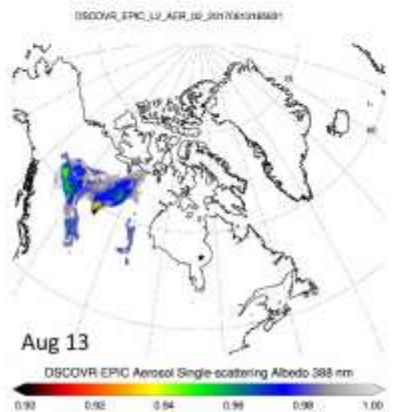
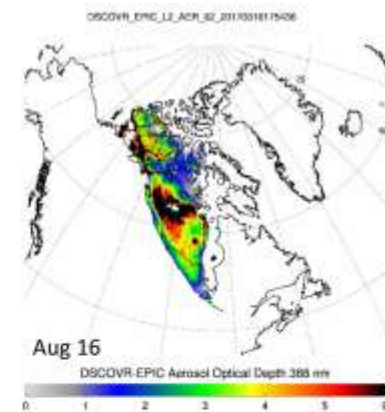
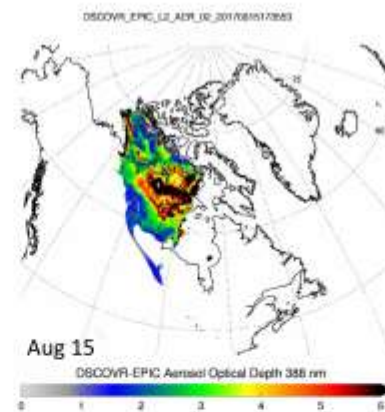
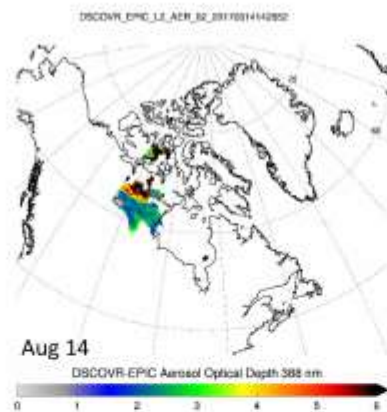
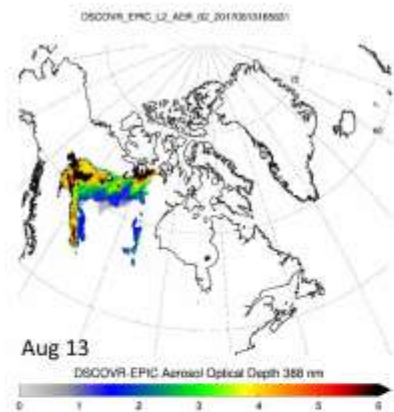
August 20, 2017



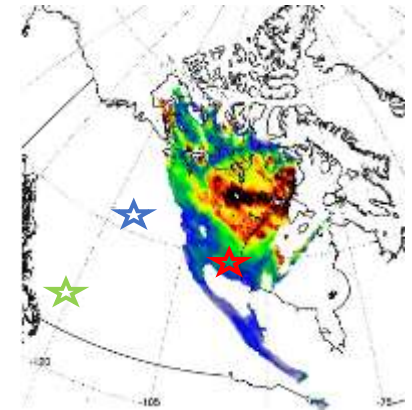
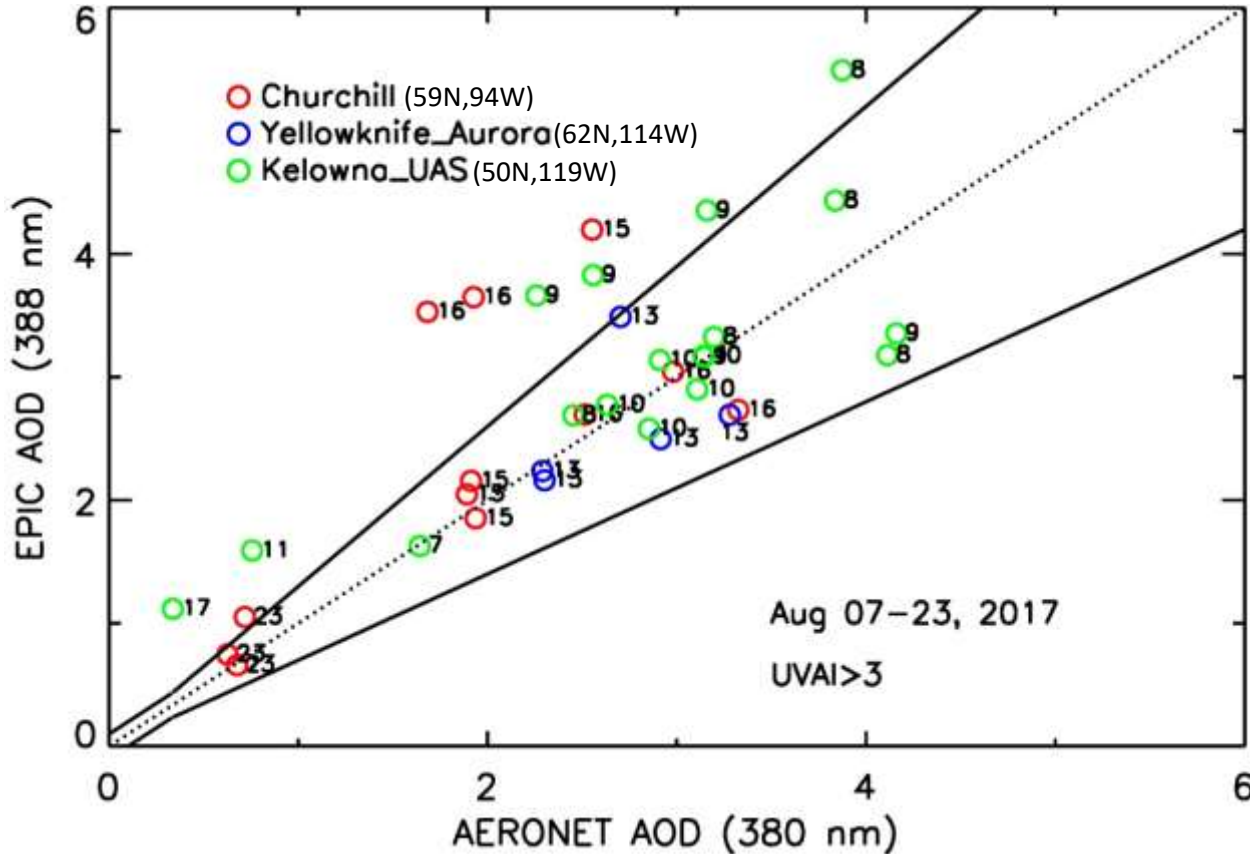
August 23, 2017



ALH constrained EPIC AOD/SSA Retrievals



EPIC-AERONET AOD comparison



Comparisons at the Kelowna_UAS site use standard EPICAERUV retrievals
Other sites use ALH (> 10 km) constrained research algorithm

Stratospheric Aerosol Mass Calculation

$$M = \Sigma \frac{4}{3} \rho r_{eff} A \tau_{str} f(r_{eff})$$

ρ is the carbonaceous aerosol particle mass density in g-cm⁻³ (0.79 -1.53)

r_{eff} is the radiatively effective radius

A is the effective geographical area associated with retrieved τ_{str}

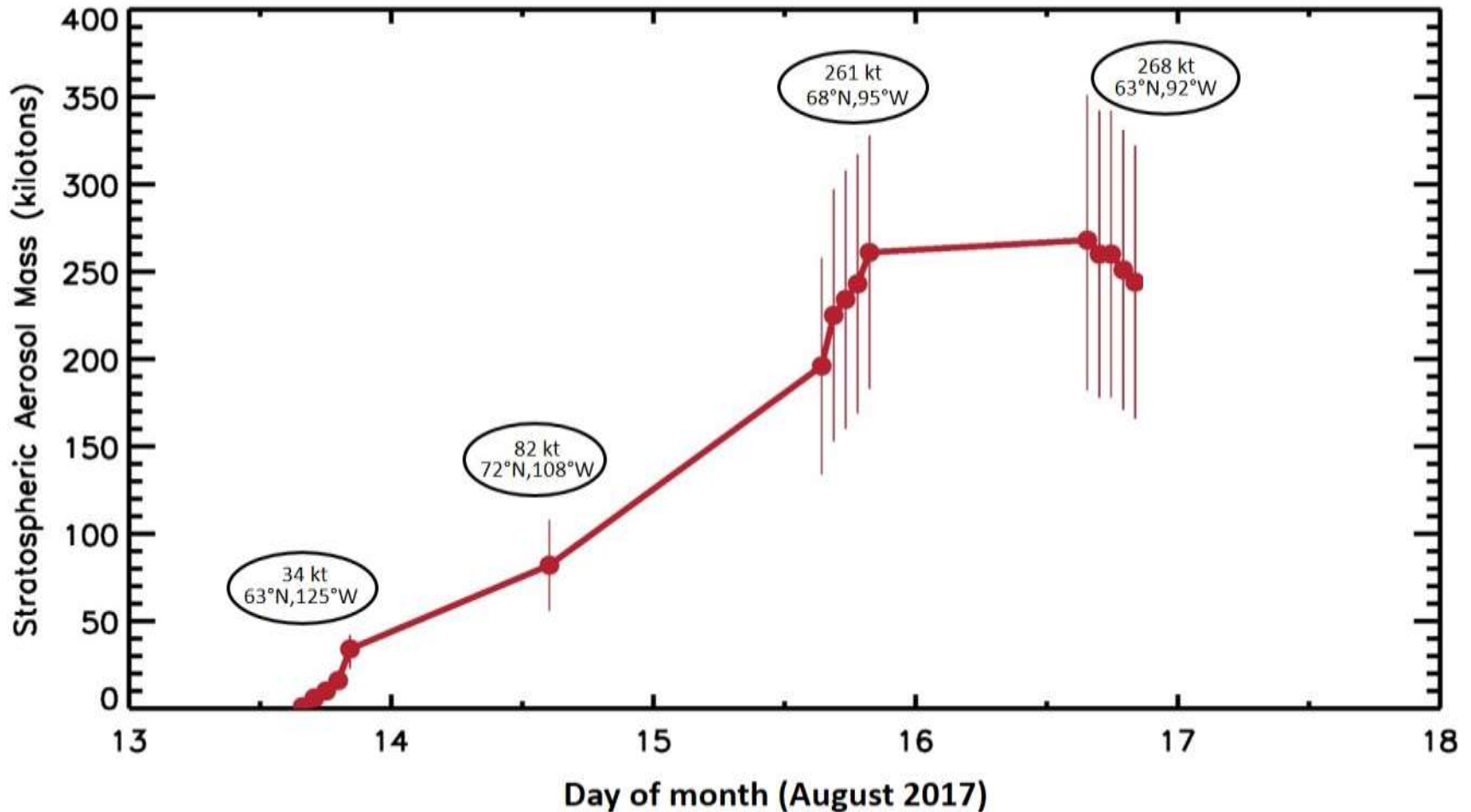
$f(r_{eff})$ is a dimensionless extinction-to-mass conversion factor, averaging over particle size distribution, defined as

$$f = \int_0^{\infty} r^2 n(r) \partial r / \int_0^{\infty} r^2 Q_{ext}(r) n(r) \partial r$$

$n(r)dr$ is the assumed number particle size distribution,

$Q_{ext}(r)$ is the extinction efficiency factor calculated using Mie scattering theory

Time line of EPIC-derived Carbonaceous Aerosol Injection in the Stratosphere



- Large day-night differences in aerosol mass injection rates are apparent.
- Observed day-night differences are consistent with solar heating processes such as production of secondary organic aerosols (photooxidation) and self-lofting resulting from aerosol absorption of solar radiation.